

Ocean Ultrasonic Shear and Compression Viscosities

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Abstract: A comprehensive description of ocean molecular flow and deformation is provided with the help of hydrodynamic and ultrasonic principles. Hydrodynamic computation of true or natural viscosities shows that ocean shear viscosity (η_G), compression viscosity (η_K), and extensional viscosity (η_E) are interrelated. There are no experimental methods available for the *in situ* measurement of these viscosities. Sound absorption coefficients (α_{obs}) allow to know the ultrasonic shear (η_{UG}), compression (η_{UK}), and longitudinal (η_L) viscosities, which decrease with increasing frequency and increase with increasing temperature, the flow activation energies having nearly equivalent values; pressure (depth) increase/decrease them at low/high frequencies. The viscosities η_{UG}^* , η_{UK}^* , η_L^* are approached at about 1000 KHz. They decrease with temperature and pressure, and increase with salinity. The η_{UG}^* becomes equal to the true shear viscosity η_G at the viscosity ratio $\delta = \eta_{UK} / \eta_{UG} = 0$.

Keywords: Chemical relaxation, ocean rheology, sound absorption, ultrasonic viscosity, viscosity ratio.

1. INTRODUCTION

The continuum mechanic approach to ocean flow and deformation (ocean rheology) has been described elsewhere [1]. It may be summarized as follows:

Turbulent diffusion prevails in the ocean over molecular diffusion. Molecular diffusion in simple fluids, as the ocean is assumed to be, is isotropic because it develops at those very small scales where stratification plays no role in inhibiting the vertical motions.

The molecular diffusion mechanism of actual seawater is provided by the physical structural relaxation of clusters [2] which contain associations with weak (water-water) and strong (ion-water) bonds with their corresponding relaxation times [3].

The molecular horizontal momentum transport of seawater is the result of three contributions: shear (η_G), extension (η_E), and compression (η_K)⁽¹⁾ flows [4] which are interrelated by Eq. (4) :

$$\eta_G = \frac{-\sigma_{12}}{\frac{\partial u}{\partial z}} = \frac{-\sigma_{12}}{\frac{1}{\Delta z} \frac{\Delta x}{\Delta t}} = \frac{-\sigma_{12}}{\dot{\gamma}_{12}} \quad (1)$$

$$\eta_E = \frac{-(\sigma_{11} - \sigma_{22})}{\frac{1}{L} \frac{\Delta L}{\Delta t}} = \frac{-(\sigma_{11} - \sigma_{22})}{\dot{\epsilon}} = \frac{-N_1}{\dot{\epsilon}} \quad (2)$$

$$\eta_K = \frac{-\Delta p}{\frac{1}{\rho} \frac{\Delta \rho}{\Delta t}} = \frac{-\Delta p}{\frac{1}{V} \frac{\Delta V}{\Delta t}} = \frac{-\Delta p}{\dot{\chi}} \quad (3)$$

$$h \eta_E + i \eta_G = j \eta_K \quad (4)$$

(1) The term compression viscosity, internationally accepted [5], is used instead of the traditional bulk viscosity.

being σ_{12} , $(\sigma_{11} - \sigma_{22}) = N_1$, and ΔP , the shear, extensional, and compression stresses, respectively ; $\dot{\gamma}_{12}$, $\dot{\epsilon}$, $\dot{\chi}$, are the shear, extensional, and compression deformation rates ; h, i, j dimensionless parameters.

The effect of the system variables on these viscosities is as follows [1]:

η_G ($1.0 - 1.5 \times 10^{-7}$ dbar.s) is independent of shear rate $\dot{\gamma}_{12}$ (Newton law, Eq. (1)), and η_K ($2 - 4 \times 10^{-7}$ dbar.s) decreases with compression rate $\dot{\chi}$ (s^{-1}).

Temperature (T) decreases η_G (Table 1) [1] and increases η_K (Table 2) [1]. The Arrhenius equation flow activation energies are positive and negative respectively: $\Delta E_G = 4$ kcal/mol, and $\Delta E_K = -1$ kcal/mol.

Table 1. Ocean True Shear Viscosities

Z (m)	P dbar	T °C	η_G^a 10^{-7} dbar.s
10	1	23.58	0.92
100	10	16.69	1.01
500	50	11.87	1.25
1000	100	7.03	1.41
1500	150	5.48	1.48
20100	200	4.03	1.54

^a Station B.

Salinity (S) increases both η_G and η_K , but almost unnoticeably.

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Decreasing pressures (P) (depth z) reduce both η_G (Table 1) and η_K (Table 2) making the clusters to move vertically from a compressed (spheroidal) to an elongated (ellipsoidal) state.

Table 2. Ocean True Compression Viscosities

Z (m)	η_K^a vs z 10 ⁻⁷ dbar.s	η_K^b vs T 10 ⁻⁷ dbar.s	T °C
10	0.20	4.12	25
100	0.40	4.06	20
500	1.20	3.99	15
1000	2.10	3.90	10
1500	3.00	3.80	5
2000	4.20	-----	---

^aT = 15°C, $\chi = 4.5 \times 10^9 \text{ s}^{-1}$, $\delta = 3.5$.

^b $\chi = 6.5 \times 10^8 \text{ s}^{-1}$, z = 250 m, $\delta = 3.5$.

There are no experimental techniques available to date to corroborate these results, but the needed supplementary experimental information may be obtained with *in situ* measurements of ocean sound absorption.

Recent developments of measurement techniques on molecular viscosity by ultrasound may be found in the literature [6-8]. Main studies on the ocean were devoted to oil slicks [9], red algae blooms [10] or the marine microlayer [11]. However to our knowledge there are no studies on the behavior of ultrasonic viscosity on the whole water column of the ocean. In section 3-a we formulate the expression of the ocean sound absorption coefficients, next in section 3-c we discuss the effect of temperature on the seawater viscosity ratio, we follow discussing the effects of pressure (depth), temperature and salinity on ultrasonic viscosities, finally we conclude summarizing the main findings.

2. EXPERIMENTAL PROCEDURES

Sound absorption coefficients (α_{abs}) were measured with the method of sofar shots detonated at different ocean locations by [12] and references there on.

Table 3. Chemical Relaxation Mechanisms

Reaction	$\overline{\Delta V^*}$ cm ³ mol ⁻¹		$\overline{\Delta \kappa^*}$ cm ³ mol ⁻¹ bar ⁻¹	
	25°C	5°C	25°C	5°C
Acid-base equilibrium (non-electrolites):				
Water $\text{H}_2\text{O} \rightarrow \text{H}^+ + \text{OH}^-$	18.0	19.5	3.15	4.73
Boric acid $\text{B(OH)}_3 + \text{H}_2\text{O} \rightarrow \text{H}^+ + \text{B(OH)}_4^-$	25.7	28.0	1.93	2.79
Carbonic acid $\text{CO}_2 + \text{H}_2\text{O} \rightarrow \text{CO}_3\text{H}_2$	22.6	24.7	1.90	3.51
$\text{CO}_3\text{H}^- \rightarrow \text{H}^+ \text{CO}_3^{2-}$	-----	-----	-----	-----
Ion-pair equilibrium:				
$\text{MgSO}_4(\text{aq}) \rightarrow \text{Mg}^{2+}(\text{aq}) + \text{SO}_4^{2-}(\text{aq})$	7.2	-----	1.7	-----

Viscosity ratios (δ) were deduced [1] from standard seawater shear viscosities [13] and ocean compressibility data [14].

3. RESULTS AND DISCUSSION

3.1. Ocean Sound Absorption Coefficients

The sound absorption coefficient (α_i) for a single relaxation process obeys the equation [15]:

$$\alpha_i = \frac{\pi \kappa_i}{\kappa_0 c} \frac{f_i f^2}{f_i^2 + f^2} \quad (5)$$

where κ_0 is the isothermal compressibility (β is the symbol commonly used for compressibility in ultrasonics), κ_i the chemical compressibility (Table 3) [16], c the sound velocity, f the acoustic frequency, and f_i the relaxation frequency which according to [17] is about 1 KHz for boric acid (B(OH)_3) (in this frequency range the contribution of scattering of inhomogeneities in the index of refraction of the medium is a constant independent of frequency, and amounts to 10⁻⁶ dB/m [18], about 10 KHz for magnesium sulfate (MgSO_4), and about 1,000 KHz for water (H_2O).

Ocean sound absorption coefficients (α_{abs} , dB/m) may be obtained from data of [12] (which are also used to assess the potential effects of suspended particles and gas bubbles) with the following equations [17]:

At salinity S = 35 [4] (A, B, C, f_1 , f_2 , being temperature, pressure, and salinity dependent parameters):

$$\alpha_{\text{obs}} = \frac{A f_1 f^2}{f_1^2 + f^2} + \frac{B f_2 f^2}{f_2^2 + f^2} + C f^2 \quad (6)$$

The first, second, and third terms of the right-hand side of Eq. (6) are related to the relaxation mechanisms of the B(OH)_3 (carbonate modify it [4]), MgSO_4 , and H_2O , respectively. See Table 3 [16] in which are also shown the partial molal volume change of the solutes in seawater ($\overline{\Delta V^*}$), and the partial molal compressibility of the solutes:

$$\overline{\Delta\kappa^*} = -\overline{\Delta V^*} / \Delta P \quad (7)$$

At $S = 35 - 37$ and frequency $f = 2 - 25$ KHz [19]:

$$\alpha_{obs} = 8.6866 \left(\frac{2.34 \times 10^{-6} S f_T f^2}{f_T^2 + f^2} + \frac{3.38 \times 10^{-6} f^2}{f_T} \right) (1 - 6.54 \times 10^{-4} P) \quad (6a)$$

with the frequencies f and $f_T = 21.9 \times 10^6 \cdot 1520 / (T + 273)$ in KHz, $T^\circ\text{C}$, P in atmosphere, and the factor 8.6866 to convert nepers/m into dB/m.

Seawater α_{obs} (Fig. 1) behaves differently at frequencies of the order of the H_2O relaxation one (500-1,000 KHz), and shows a change of slope at about 15°C . The molecular

structure of pure water is [20] (Fig. 1c): hexagonal (ice crystals) below 0°C , tetrahedral (pentamer) above 5°C , chain like (trimer) at about 27°C , more than 10% of the water molecules have changed their tetrahedral structure to the chain one above 30°C .

The ocean sound absorption coefficient is related to the ultrasonic shear (η_{UG}) and compression (η_{UK}) viscosities by the equation [17]:

$$\alpha_{theor} = \frac{16 \pi^2 f^2}{3 \rho c^3} \left(\eta_{UG} + \frac{3}{4} \eta_{UK} \right) = \frac{16 \pi^2 f^2}{3 \rho c^3} \eta_{UG} + \frac{4 \pi^2 f^2}{\rho c^3} \eta_{UK} \quad (8)$$

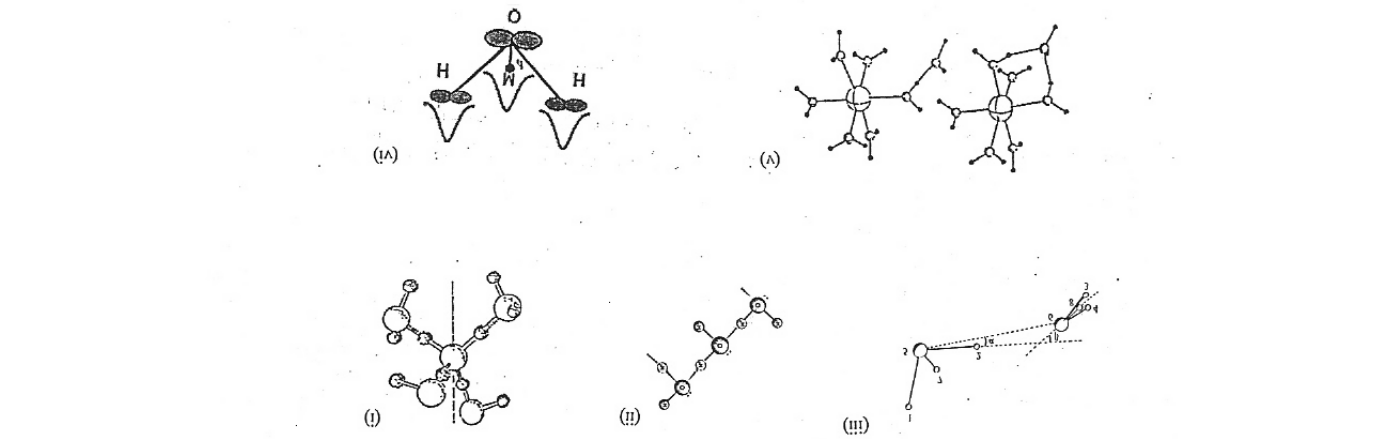
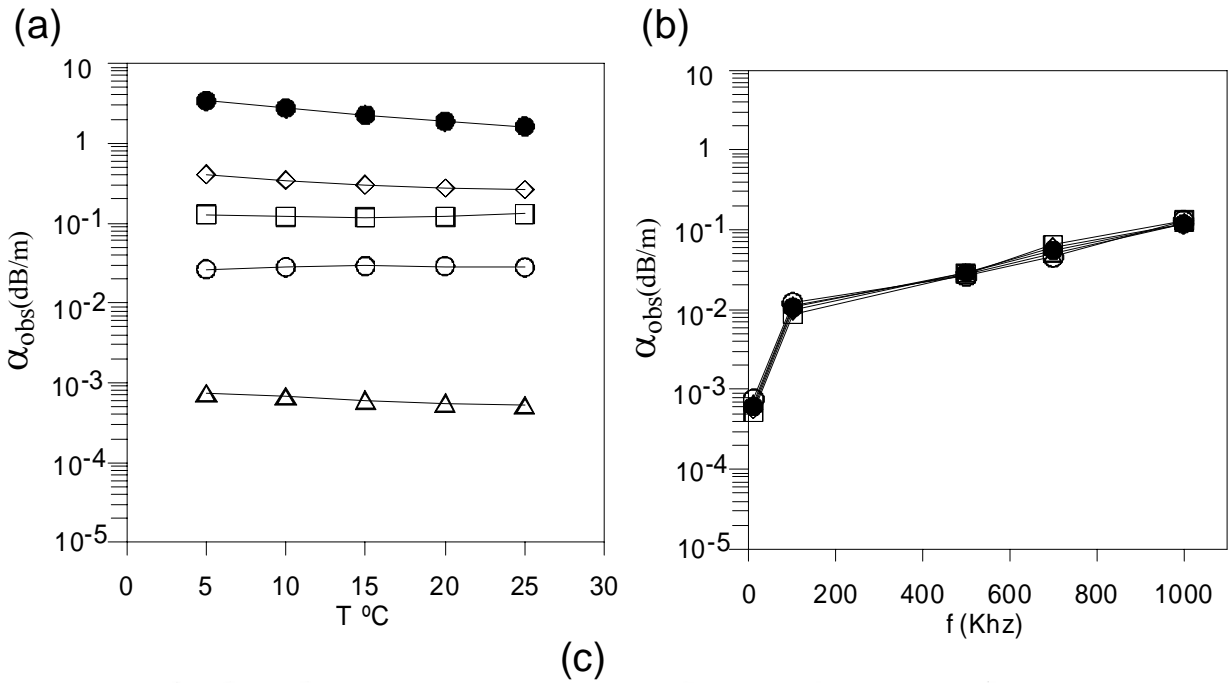


Fig. (1). Ocean sound absorption coefficient (α_{obs}) at $S = 35$, depth $z = 1,000$ m, $\delta = 3.2$, (a) as a function of temperature at (Δ) 10 KHz, (\circ) 100 KHz, (\square) 500 KHz, (\diamond) 1,000 KHz, (\bullet) 3,000 KHz. (b) as a function of frequency at (\circ) 5°C , (Δ) 10°C , (\bullet) 15°C , (\diamond) 20°C , (\square) 25°C . (c) Pure water structure: (i) pentamer (\bullet hydrogen atom, \circ oxygen atom, from [24]), (ii) trimer (from [10].), (iii) dimer (from [25]), (iv) monomer (from [26]), (v) ion cluster illustrating one water molecule in the second layer (from [2]).

in which f is the frequency in Hz, and the viscosities expressed in poises, the density in g/cm^3 , and the sound velocity c in cm/s calculated with the algorithm described by [17].

The first term of Eq. (8) is the absorption coefficient α_{class} in the low frequency limit, i.e., the shear viscous loss [4]:

$$\alpha_{class} = \frac{16\pi^2 f^2}{3\rho c^3} \eta_{UG} \quad (9)$$

3.2. Viscosity Ratio

The shear and compression viscosities are interrelated by the equation [1]:

$$\delta = \eta_{UK} / \eta_{UG} \quad (10)$$

Measurement of this dimensionless ratio provide $\delta = 2.81$ for seawater [18]. Seawater δ values at $S = 35 - 37$, $T = 5^\circ\text{C} - 25^\circ\text{C}$, $P = 1 - 2000$ dbar, may be deduced as in [1].

Fig. (2a) is a typical example at $S = 35$ since: i) the η_G used are for standard seawater [13], ii) the effect of S on both η_G and η_K are almost unnoticeable [1], iii) Eq. (6) used to compute the sound absorption coefficients (α_{obs}) is valid only for $S=35$ [17]. As may be observed the plots are nearly straight lines and δ increases with increasing depth (z) and decreasing compression rate ($\dot{\chi}$).

The effect of temperature on the seawater viscosity ratio (Fig. 2b) shows a different slope below and above 15°C [1], what possibly is due to the fact that seawater is a dilute polymeric system consisting of 6-60 water molecules with an ion in the center (solvation shell) (Fig. 1c). Under the pressure of sound waves the relaxation mechanism is a chemical association-dissociation reaction [15] giving rise to a large

volume change upon ionization (Table 3, [16]). Most seawater cations are in a 99% dissociated state, while magnesium sulfate (MgSO_4) is only 78.5% dissociated (Table 4) [21]. As the temperature increases a decrease of both the partial molal volume change ($\overline{\Delta V^*}$) and the partial molal compressibility ($\overline{\Delta \kappa^*}$) occurs (Table 3). The corresponding cluster structure expands decreasing the volume deformation ($\chi = (\Delta V / V)$) [1]. The final result is an increase in compression viscosity and viscosity ratio (Eqs. (3) and 10)).

The viscosity ratio δ of pure water has been measured [4] to be 2.68 -2.75 and is almost independent of depth (pressure), while it decreases slightly at all temperatures [4, 22] and, at frequencies of the order of GHz, decreases steeply below 15°C (Figure 8 of [23]); compare with Fig. (1a) and Eq. (11).

3.3. Ultrasonic Viscosities

The ultrasonic compression viscosity (η_{UK}) is [4]:

$$\eta_{UK} = \frac{4}{3} \eta_{UG} \left(\frac{\alpha_{obs} - \alpha_{class}}{\alpha_{class}} \right) \quad (11)$$

Substituting in Eq.(11) the Eqs.(6-6a, 9, 10) results:

$$\eta_{UK} = \delta \cdot \eta_{UG} = \frac{4}{3} \eta_{UG} \left(\frac{\alpha_{obs}}{\alpha_{class}} - 1 \right) \quad (12)$$

which allows to find the ultrasonic shear viscosity (η_{UG}) as follows (the factor 0.25 comes from $0.25 = (4/3) / (16/3)$):

$$\delta \cdot \eta_{UG} = \frac{4}{3} \eta_{UG} \cdot \frac{\alpha_{obs}}{\frac{16\pi^2 f^2}{3\rho c^3} \eta_{UG}} - \frac{4}{3} \eta_{UG} = \frac{0.25 \cdot \alpha_{obs} \cdot \rho c^3}{\pi^2 f^2} - \frac{4}{3} \eta_{UG} \quad (13)$$

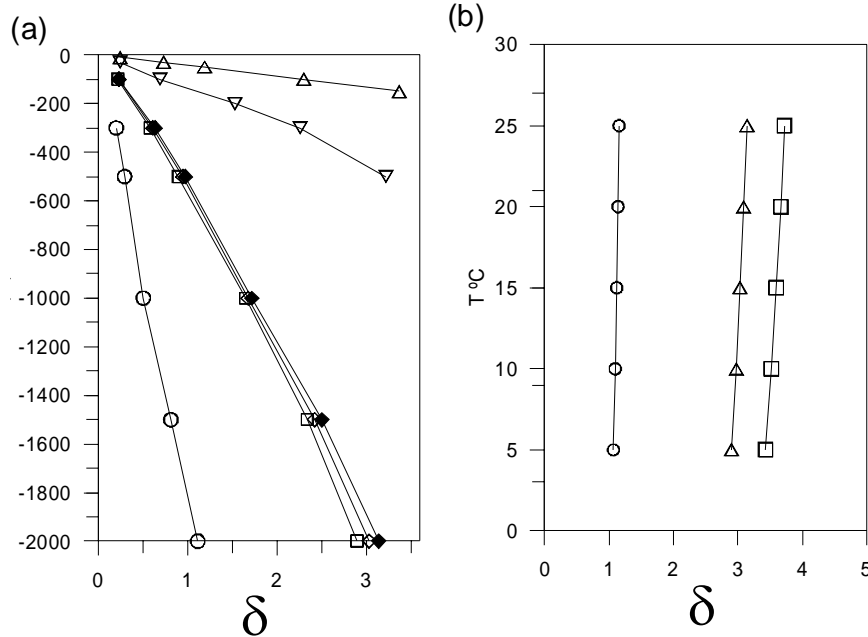


Fig. (2). Seawater viscosity ratio (δ) (a) at $T = 15^\circ\text{C}$, $S = 35$, as a function of depth (z) and different compression rates ($\dot{\chi}$): (Δ) $4.2 \times 10^8 \text{ s}^{-1}$, (∇) $1.2 \times 10^9 \text{ s}^{-1}$, (\diamond) $4.25 \times 10^9 \text{ s}^{-1}$, (\blacklozenge) 25°C , (\square) 5°C , (\circ) $1.1 \times 10^{10} \text{ s}^{-1}$. (b) as a function of temperature at $S=35$, $z = 2,000$ m, (Δ) $\dot{\chi} = 4.2 \times 10^9 \text{ s}^{-1}$, (\circ) $\dot{\chi} = 1.1 \times 10^{10} \text{ s}^{-1}$, (\square) $1.16 \times 10^9 \text{ s}^{-1}$ ($z = 500$ m).

Table 4. Dissociation States of Seawater Salts

Free Ions			% Associated With						
Cation	g/kg (15)	%					SO ₄ ²⁺	CO ₃ H ⁻	
Sodium	10.733	99					1.2	0.01	
Magnesium	1.294	87					11.0	1.0	
Calcium	0.412	91					8.0	1.0	
Potassium	0.339	99							
Strontium	0.0079								
Anion			Na ⁺	Mg ²⁺	Ca ²⁺	K ⁺			
Chloride	19.354		----- reference ion-----						
Sulfate	2.712	54	21	21.5	2	0.5			
Bicarbonate	0.142	69	8	19.0	4	-----			
Carbonate	-----	9	17	67.0	4				
Borate	0.0045	0.01							
Bromide	0.067								
Fluoride	0.0013								

Therefore

$$\delta \cdot \eta_{UG} + \frac{4}{3} \eta_{UG} = \left(\delta + \frac{4}{3} \right) \eta_{UG} = \frac{0.25 \cdot \alpha_{obs} \cdot \rho c^3}{\pi^2 f^2} \quad (14)$$

and (using the factor 10⁻² to have the results with units 10⁻⁷ dbar.s) the ultrasonic shear viscosity is:

$$\eta_{UG} = \frac{0.25 \cdot 10^{-2} \cdot \alpha_{obs} \cdot \rho c^3}{\pi^2 f^2} \cdot \frac{1}{\delta + \frac{4}{3}} \quad (15)$$

which at δ = 0 is equivalent to Eq. (9) (see also Fig. 3d).

The ultrasonic compression viscosity (η_{UK}) is obtained then from Eq. (12):

$$\eta_{UK} = \delta \cdot \eta_{UG} = \frac{0.25 \cdot 10^{-2} \cdot \alpha_{obs} \cdot \rho c^3}{\pi^2 f^2} \cdot \frac{\delta}{\delta + \frac{4}{3}} \quad (16)$$

which at δ = 0 becomes η_{UK} = 0 (Fig. 3f).

The longitudinal viscosity (η_L) may be estimated as [18]:

$$\eta_L = \frac{4}{3} \eta_{UG} + \eta_{UK} \quad (17)$$

Fig. (3a, c, e, g), show that the three viscosities η_{UG}, η_{UK}, η_L decrease with increasing frequency, becoming independent of it at about 20 KHz (Fig. 3b). The viscosities η^{*}_{UG}, η^{*}_{UK}, η^{*}_L, shown in Fig. (3d, f, h), are approached at the H₂O relaxation frequency (1,000 KHz), as shown in Fig. (3c, e, g).

As δ decreases:

η^{*}_{UG} increases towards the true viscosity η_G (Table 1) [1].

η^{*}_{UK} decreases towards η^{*}_{UK} = 0. It also decreases with temperature: note that this is contrary to what happens with continuum mechanics η_K values (Eq. 3) which increase with temperature (Table 2) [1]. It is due to fact that the effect of temperature is different on the chemical association-dissociation relaxation than on the physical cluster structure relaxation. The activationn energies (ΔE) computed with Arrhenius equation [1] have nearly equivalent values (shear ΔE_{UG} = 2.46 kcal/mol, compression ΔE_{UK} = 1.80 kcal/mol, longitudinal ΔE_L = 2.10 kcal/mol).

η^{*}_L is almost constant suggesting that (Eqs. (17) and (4)):

$$\eta_L^{*(\delta \rightarrow 0)} = \frac{4}{3} \eta_{UG}^{*(\delta \rightarrow 0)} + \eta_{UK}^{*(\delta \rightarrow 0)} \approx h \eta_E = -i \eta_G + j \eta_K \quad (18)$$

It shows the effect on η_E of the strength of the cluster hydrogen bonds (which are, for example, stronger for Na⁺ and weaker for Cl⁻), and that the clusters are more compacted at large depths.

Fig. (4a) shows that as the temperature increases η_{UG}, η_{UK}, η_L increase at low frequencies up to 15°C and decrease at higher temperatures for all frequencies. It is the result of the combined effects of the sound velocity c (increase with increasing temperature [17]), the sound absorption coefficient α_{obs} (Fig. 1), and the viscosity ratio δ (Fig. 2b). η^{*}_{UG}, η^{*}_{UK}, η^{*}_L (Fig. 4b) decreases with temperature.

The effect of pressure (depth) is to increase η_{UG}, η_{UK}, and η_L (Fig. 4c). The viscosities η^{*}_{UG}, η^{*}_{UK}, and η^{*}_L above 1,000 KHz (Fig. 4d) decrease since increasing depth (P) decrease Δκ^{*} (equation 7).

The effect of salinity (S) on the ultrasonic viscosities can be known replacing Eq. (6) by Eq. (6a) in Eq. (15). Fig. (4e, f) show that the salinity increases the three viscosities since increasing salinity decreases ΔV^{*} and Δκ^{*}.

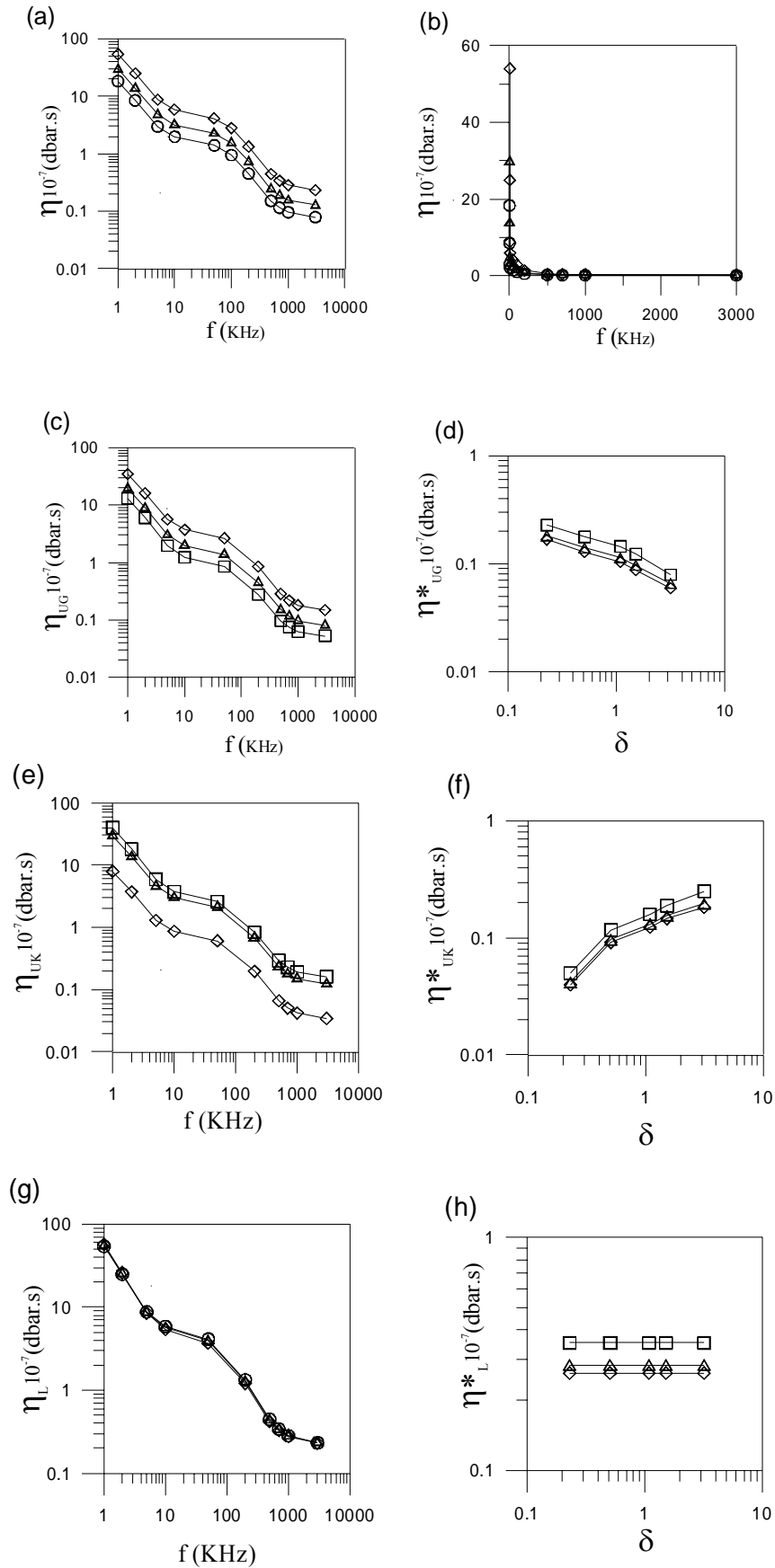


Fig. (3). Ultrasonic viscosities at $T = 15^{\circ}\text{C}$, $S = 35$, $\delta = 1.52$, as a function of frequency: **(a)** η_{UG} , **(b)** η_{UG}^* , **(c)** η_{UG} , **(d)** η_{UG}^* , **(e)** η_{UK} , **(f)** η_{UK}^* , **(g)** η_L , **(h)** η_L^* , and at different viscosity ratios δ (\square) 80.23, (Δ) 1.52, (\diamond) 3.03, **(c)** η_{UG} , **(e)** η_{UK}^* , **(g)** η_L^* . As a function of δ at frequency $f_i = 1,000$ KHz and different temperatures (\square) 5°C , (Δ) 15°C , (\diamond) 25°C); **(d)** η_{UG}^* , **(f)** η_{UK}^* , **(h)** η_L^* .

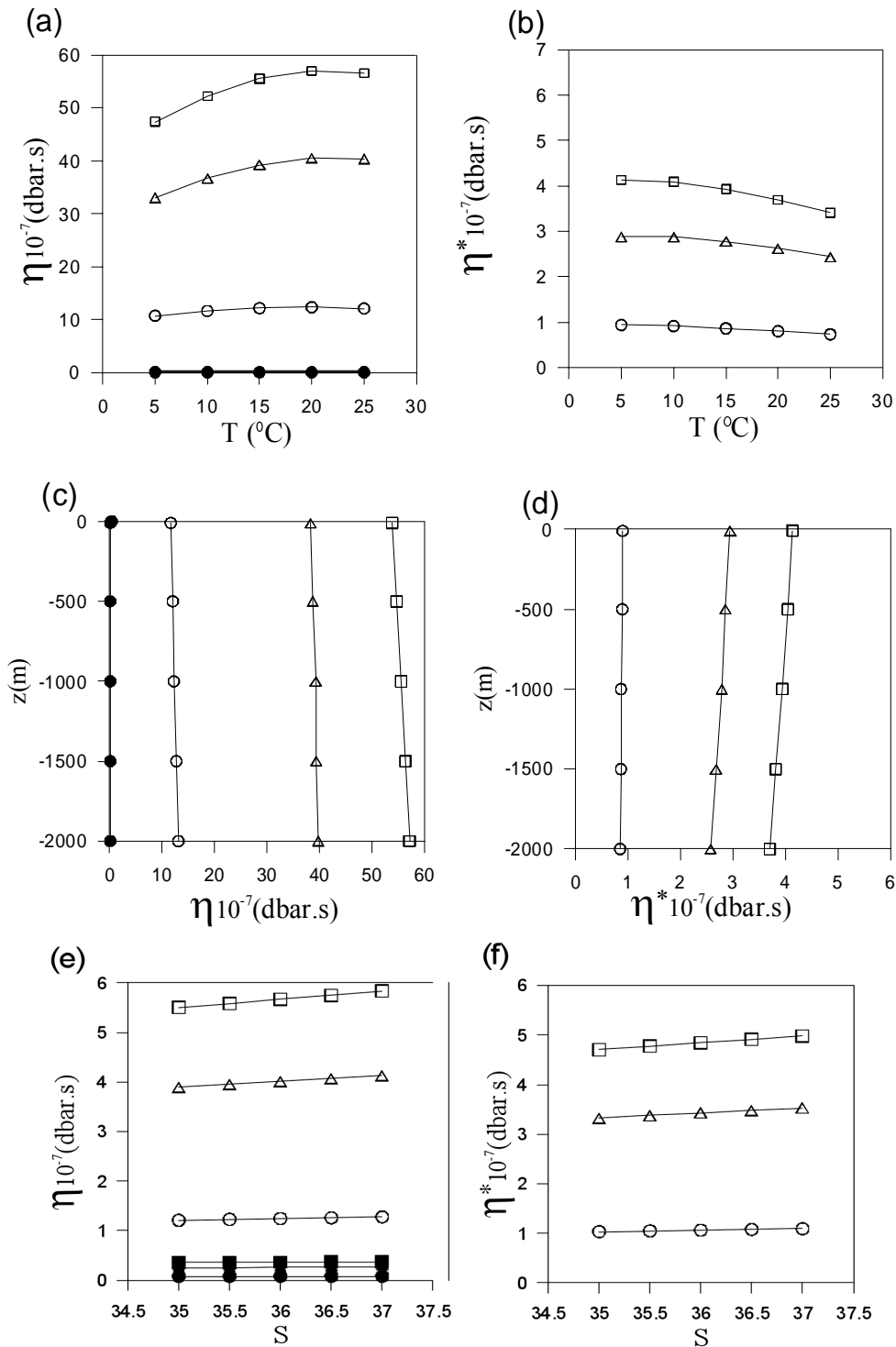


Fig. (4). Effect of temperature at $S = 35$, $z = 1,000$ m, $\delta = 3.2$, on (a) (\circ) η_{UG} , (Δ) η_{UK} , (\square) η_L , at frequency (white symbols) 1KHz, and (\bullet) 700 KHz. (b) (\circ) η^*_{UG} , (Δ) η^*_{UK} , (\square) η^*_L ($f_i = 1,000$ KHz). Effect of depth (z) at $T = 15^\circ\text{C}$, $S = 35$, $\delta = 3.2$, and frequency (white symbols) 1 KHz, and (\bullet) 700 KHz: on (c) (\circ) η_{UG} , (Δ) η_{UK} , (\square) η_L (\bullet , \square , at 700 KHz); (d) (\circ) η^*_{UG} , (Δ) η^*_{UK} , (\square) η^*_L ($f_i = 1,000$ KHz). Effect of salinity (S) at $T = 15^\circ\text{C}$, $z = 1,000$ m, $\delta = 3.21$, and $f = 10$ KHz: (e) (\circ) η_{UG} , (\diamond) η_{UK} , (\square) η_L (\bullet , \blacktriangledown , \blacksquare , at 700 KHz); (f) at $f_i = 1,000$ KHz (\circ) η^*_{UG} , (Δ) η^*_{UK} , (\square) η^*_L .

4. CONCLUSIONS

Seawater is a system which contains compressed hydrogen-bonded reversible ion networks (clusters) suspended in water oligomers. The molecular behavior of this system may be interpreted with two different

approaches. The continuum mechanics with which the shear, extensional, and compression viscosities may be evaluated [1]. Ultrasonic measurements which show that ultrasonic shear, ultrasonic compression, and longitudinal viscosities decrease with frequency, increase/decrease with temperature,

and the pressure (depth) increase/decrease them at low and high frequencies, respectively. At the lower viscosity ratio (δ) the ultrasonic shear viscosity is similar to the continuum mechanics one.

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