

A Comparison of Equatorial Electrojet in Peru and East Brazil

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Abstract: The paper describes the characteristics of the equatorial electrojet at Huancayo (HUA, 12.1°S, 75.3°W, inclination 1.5°N, declination 1.0°E) in western side of South America, where the geomagnetic field is aligned almost along the geographic meridian, and at Itinga (ITI, 4.3°S, 47.0°W, inclination 1.4°N, declination 19.3°W) in eastern part of South America, where the geomagnetic field is aligned about 19° west of the geographic meridian; although the mean intensity of the magnetic field in the two regions are almost of the same order. Further comparisons are made of the current at Itinga and at Tatuoca (TTB, 1.2°S, 48.5°W, inclination 7.8° N, declination 18.7°W), a low latitude station in the same longitude sector. The daily range of horizontal component of the geomagnetic field, H, is shown to be almost 16% higher at HUA compared to that at ITI. The daily variation of the eastward field, Y, showed a strong minimum of -40 nT around 13-14 hr LT at ITI whereas very low values were observed at HUA with a positive peak of about 4 nT around 11-12 hr LT. The vertical field, Z, showed abnormally large negative values of -70 nT at TTB around 13 hr LT. The day-to-day fluctuations of midday and midnight values of X field were positively correlated between HUA and ITI with a high correlation coefficient of 0.78 and 0.88 respectively. Values of Y field were also significantly positively correlated between HUA and ITI for midnight hours (0.72), while no correlation was observed for the midday hours. The midnight values of X field at HUA, ITI and TTB showed significant (0.90 or greater) correlation with Dst index. Correlation values of about 0.7 were observed between Dst and midday values of X at ITI and TTB and to a lesser degree (0.4) at HUA.

Keywords: Equatorial electrojet, equatorial E-region, geomagnetism.

1. INTRODUCTION

The most important discovery which has revolutionized the basic idea of low latitude ionospheric current system had been the result of the ground magnetic survey by Giecke (only quoted by Chapman [1]) at fourteen stations in Peru during the period September to November 1949. The daily range of the horizontal geomagnetic field, H, was found to increase from 7° geographic latitudes to a peak at 13°S (close to the magnetic dip equator) with a ratio of 2. Chapman [1] interpreted the result as due to a narrow band of thin current sheet flowing eastward during day time hours in the ionosphere and named as equatorial electrojet (EEJ).

This sudden increase in the daily range of H near magnetic equator (Huancayo anomaly) presented a serious problem to the Chapman [2] current system which was derived from data at only middle latitudes, now requiring abnormally large local wind velocities to generate a large electro motive force for the EEJ. This problem was solved by Baker and Martyn [3] who pointed that near the magnetic equator the geomagnetic field, electric field and the vertical electron density gradient are orthogonal to each other. At the magnetic equator the direction of current flow during the day

time is eastward and the magnetic field is northward. The electrons within 70-140 km altitude drift upward relative to the ions causing a vertical Hall polarization field, which increases the eastward Cowling conductivity. They also explained that beyond 3° north or south from the magnetic equator, the polarization field leaks away along the inclined magnetic field and thus the abnormal conductivity is restricted to ±3° dip latitudes. Thus, the Sq electric field extended from low to equatorial latitudes was shown to be sufficient to generate the large equatorial electrojet current. Equatorial electrojet has been extensively studied from ground, rocket and satellite based magnetometers, spaced receiver drift, VHF backscatter radar and theoretical studies [4-6].

Rastogi [7] showed that the magnitude of EEJ is most pronounced in American zone and least in Indian zone. He suggested an inverse relationship between the equatorial conductivity and the background magnetic field in the environment. Model calculations have been made of EEJ [8-11]. Richmond [12] developed a numerical model including winds and instabilities and showed that the two-stream instability limits the strength of the polarization electric field and the eastward current. Increase in magnetic field strength reduces the Pedersen and Hall conductivities at all levels and also lowers the altitude of maximum polarization field to a level where Hall conductivity is lower thus reducing the electrojet strength. Gangepain *et al.* [13] compared different models and showed that differences arise largely due to

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different values of electron collision frequencies used. By empirically adjusting the electron collision frequency it was shown that a value that is four times that calculated from laboratory data correctly modeled the electrojet parameters. Ronchi *et al.* [14] examined how fully developed strongly turbulent state of the small-scale waves (tens of meters size) affects the large-scale dynamics. Adding the additional mobility and diffusion and the small wave-length density fluctuations from in-situ rocket measurements made from Peru it was shown that the equilibrium vertical electric field peaks around 100 km in absence of turbulent terms and around 105 km in presence of them. Peak value of the vertical polarization field also decreases in the presence of turbulent terms.

Ionospheric drift measurements from spaced receiver technique at Thumba showed a high correlation between drift velocity and electrojet strength as determined from the difference between the ranges in H at Trivandrum, near dip equator and Alibag, outside electrojet [15]. Thus in the absence of electric field measurements one can use electrojet strength as an index of the electric field in ionosphere (E-region). Anderson *et al.* [16] showed quantitative relationship based on the Jicamarca incoherent scatter radar drift observations and magnetometer data at Canete and Piura in Peru. It was further demonstrated for the Philippine longitude sector [17]. Uozumi *et al.* [18] have studied EEJ using MAGDAS/CPMN data at four stations Addis Ababa (AAB), Davao (DAV), Ancon (ANC) and Eusebio (EUS) and found that the mean night time (18-06 LT) value of the H shows variations like Dst and can be used as a proxy to Dst for real-time and long-term monitoring and defined it as EDst index.

It may be mentioned here that the role of abnormally large wind was shown to affect the daily variations of H at Huancayo on certain days by Bartels and Johnston [19,20]. They showed that the lunar variations in H were abnormally large at Huancayo, such that on certain days the ΔH during the daytime hours decreases below the base nighttime values. Gouin and Mayaud [21] described the phenomenon as observed at Addis-Ababa and named it as counter electrojet.

Day-to-day variability of SqH at mid latitude has been described by Greener and Schlapp [22] who found the longitudinal spatial scale to be of the order of 2000 km. James and Rastogi [23] have described day-to-day fluctuations in H field at 13 stations along Indo-Russian longitude sector and found regions of high correlation among themselves: one equatorial region, second low latitudes equatorward of Sq focus and the third at mid latitudes north of Sq focus. Rastogi *et al.* [24] studied the day-to-day fluctuations in the daily range of H at the EEJ stations, Ancon (ANC) in Peru (77.6 W) and at Sao Luis (SLZ, 44.2° W) during the period January–June 1993. Regardless of the month, the range of H was significantly larger at ANC than at SLZ. There was excellent correlation in the daily range of H at the two stations.

Features of EEJ for different longitude sectors have been described for 75°W by [25]; for 15°E [26]; for 75°E [27, 28]; for 60°W [29] and for 45°W [30-33]. Rastogi and Trivedi [31] have described the profile of EEJ current system in the East Brazilian region. The most important feature of the

geomagnetic field in this region has been that the horizontal field vector deviates by about 20° westward of the geographic north. Thus very large westward component of the field was observed compared to geomagnetic field in the western part of South America where it is directed almost towards the geographic north. Kane and Trivedi [34, 35] compared equatorial electrojet at Huancayo in Peru and Eusebio in east Brazil. Shume *et al.* [36] have recently reported seasonal and longitudinal variability of the equatorial electrojet at Jicamarca and Sao Luis (East Brazil) for both solar maximum (2001-2002) and solar minimum (2007-2007).

2. DATA

A unique dense network of 26 three-axis fluxgate magnetometers were operated in East Brazil within the latitude 7.8°S to 4°N and longitudes 35°W to 49°W from November 1990 to March 1991. Digital fluxgate magnetometers measured vertical (Z) and two orthogonal horizontal components of the geomagnetic field every minute with a resolution of 1 nT. The horizontal components were resolved to geographic north (X) and east (Y) using total horizontal field and the magnetic declination values of the stations. The maximum temperature changes of the fluxgate sensors were only 2-3° during the period of observations so the temperature effects are negligible. Rigoti *et al.* [30] have described the characteristic of the EEJ from these data.

In this paper we present the features of the equatorial and low latitude current system in the eastern and western regions of South America based on comparison of the data from Huancayo, HUA (12.1°S, 75.3°W, inclination 1.5°N, declination 1.0°E) with from Itinga, ITI (4.3°S, 47.0°W, inclination 1.4°N, declination 19.3°W). Further to study the current system outside the EEJ we have compared data from ITI with that at Tatuoca, TTB (1.2°S, 48.5°W, inclination 7.8° N, declination 18.7°W). The station code, geographic coordinates, geomagnetic inclination and declination angles at the locations are listed in Table 1. It is to be noted that the background geomagnetic field at both EEJ stations are almost same being about 28000 nT but declination is 1.0°E at Huancayo and 19.3° W at ITI. The distance between the two stations is 28° in longitude (1.9 hours) or about 2000 km.

Table 1. Geographic Coordinates and Magnetic Parameters of the Stations

Name	Code	Geog. Lat	Geog. Long	Inclination	Declination
Tatuoca	TTB	1.2 S	48.5 W	7.8 N	18.7 W
Itinga	ITI	4.3 S	47.6 W	1.4 N	19.3 W
Huancayo	HUA	12.1 S	75.3 W	1.5 N	1.0 E

We define various terms used in the analyses. ΔX , ΔY and ΔZ are the deviations of the hourly mean at the local time from the corresponding values between 0000 and 0100 hr LT (with very low value of conductivity in night time one does not expect ionospheric currents). The direction of H vector $\theta^0 = \arctan \Delta Y/\Delta X$ is measured positive east of geographic north. The disturbance ring current index, Dst

have been downloaded from the website of WDC for geomagnetism in Kyoto. The period of study chosen is 23 November 1990 to 31 December 1990 for which the data were available at all the three stations. The average sunspot number during this period was 142 and the average Ap index was 8. The Dst index were very low for most of the days except during the sudden commencement storm starting at 2330 UT on 26 November 1990 when its value dropped to more than -100 nT. The average values therefore are not truly representative of the international quiet days but fairly close to the geomagnetic quiet conditions.

Rastogi *et al.* [33] made quantitative estimate of the standard error in mean $\frac{\sigma}{\sqrt{n}}$ and of the standard deviation (σ)

of ΔX , ΔY and ΔZ at each hour at SLZ for the month of April 1993. The standard error in mean in ΔX varied from around 8 nT at noon to 5 nT in the evening-midnight hours. The value of the standard error in mean in ΔX was least in the midnight-morning hours (1-2 nT). The standard deviation was 40 nT around noon and remained fairly high in the evening-midnight hours with values around 30 nT. The standard error in mean ΔY was maximum around noon with a value of about 5 nT. The standard deviation was also maximum around noon with a value of 23 nT and decreased to 7-8 nT in the evening-midnight hours. The high values of the standard deviations at noon are partly because of the day-to-day variability in the electrojet. Based on the mean values and standard errors in mean of ΔX and ΔY at SLZ, the errors in the direction of H vector were computed. Error was very small (about a degree) between 09h to 15h LT. The error is very high between sunset and sunrise hours due to low values of the deviations in ΔX and ΔY because of very low value of conductivity. For the present data set of duration of about a month the errors will be of similar order.

3. RESULTS AND DISCUSSION

3.1. Mean Daily Variations of ΔX , ΔY , ΔZ and θ

The mean daily variations of ΔX , ΔY , ΔZ and θ at the three stations are shown in Fig. (1a, b). The ΔX was maximum of 128 nT around 11 h LT at HUA, 104 nT at 12 h LT at ITI and 66 nT at 12 h LT at TTB. The value of ΔH obtained from ΔX and declination was 110 nT at ITI and thus it is important to note that the range of H is about 16% higher at Huancayo than at Itinga, inspite of the main magnetic field (H) same at the two stations. Stronger electrojet currents in western region than the eastern region of South America may arise from stronger electric field and/or higher conductivities. The range of H at Tatuoca was 60 nT compared to 104 nT at Itinga, because it is outside the EEJ belt. Kane and Trivedi [34-35] compared the equatorial electrojet at Huancayo in Peru and Eusebio in east Brazil and found electrojet stronger at Huancayo than at Eusebio. Recently Shume *et al.* [36] have reported seasonal and longitudinal variability in the equatorial electrojet at Jicamarca in Peru and Sao Luis in east Brazil during solar maximum (20001-2002) and solar minimum (2006-2007). The equatorial electrojet was stronger in the west coast than in the east coast for both solar maximum and solar

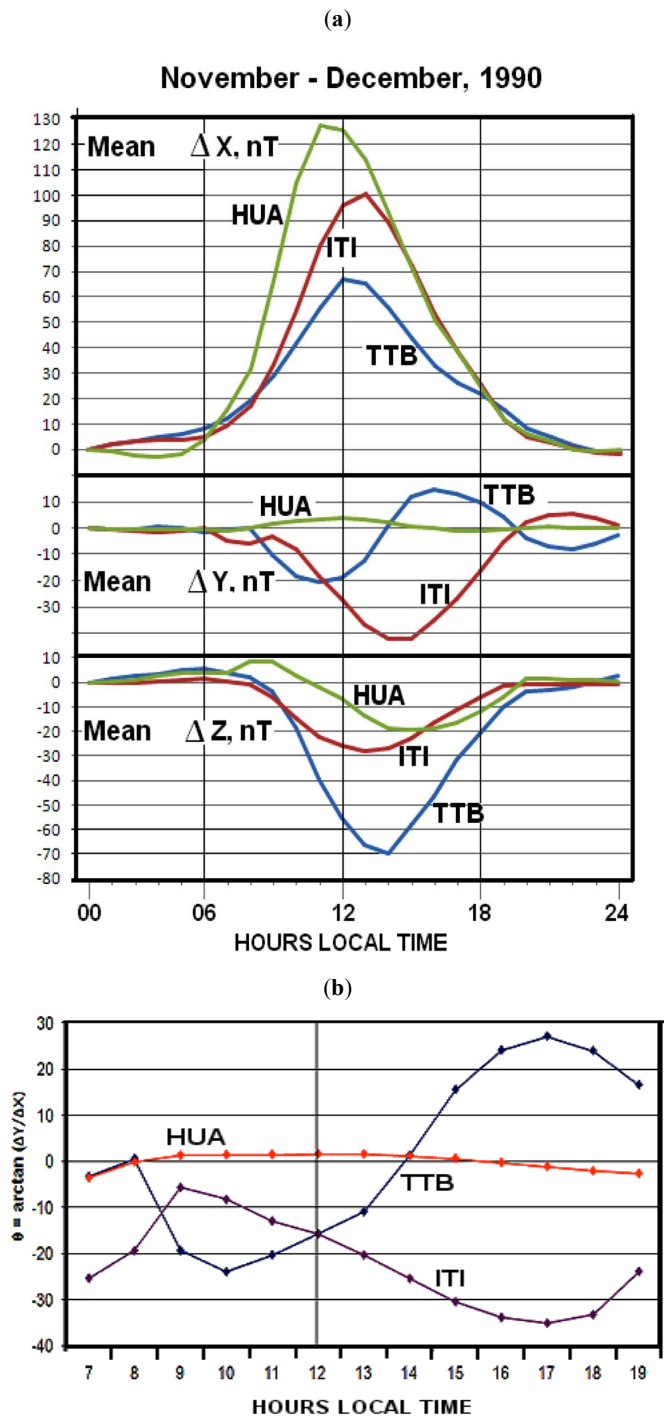


Fig. (1). Daily mean variations of (a) the deviations in X, Y and Z components at HUA, ITI and TTB and (b) the direction of the horizontal vector H over HUA, ITI and TTB averaged over the period 22 November – 31 December 1990.

minimum. Seasonally while equatorial electrojet shows equinoctial peaks at Jicamarca there is maximum during December-January at Sao Luis. The observed seasonal difference was suggested to be likely due to the large declination at Sao Luis. It was also shown that the magnitude of the seasonal and solar cycle variability is more pronounced at Sao Luis than at Jicamarca.

The value of ΔY showed a peak of about 4 nT at 11 h LT at HUA but minimum value of -40 nT at 13-14 h LT at ITI.

ΔY at TTB showed a minimum of -20 nT in the forenoon and maximum of +18 nT in the afternoon, a characteristic of December solstices [37].

The daily variation of ΔZ showed a large midday afternoon minimum at each of the stations. The magnitude of ΔZ at TTB was very large as the station lies close to the northern fringe of the EEJ belt, in accordance to the Chapman's [1] model of the equatorial electrojet.

The direction of H vector (θ) at HUA was about -3° after sunrise, rose to $+2^\circ$ around noon and decreased to -2° by the evening hours. The direction of H vector at ITI was negative throughout the day varying from -25° in the morning to -6° at 08h LT and thereafter steadily reaching to -35° at 15h. The direction of the H vector at TTB was about -20° in the forenoon and $+30^\circ$ in the afternoon hours. There may be some contribution of field-aligned currents in ΔY .

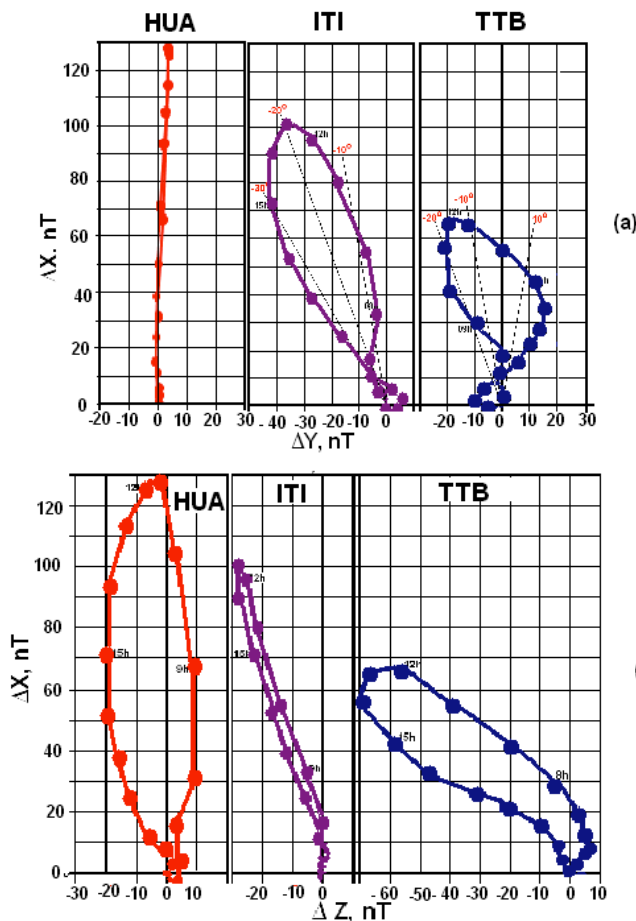


Fig. (2). (a) Hourly mean values of the deviations ΔX plotted as a function of ΔY over HUA, ITI and TTB, (b) Hourly mean values of the deviations ΔX plotted as a function of ΔZ over HUA, ITI and TTB.

In Fig. (2a) are drawn the loops showing the simultaneous diurnal development of ΔY and ΔX at HUA, ITI and TTB. The dashed lines indicate directions of regression lines (not shown for HUA where points lie along the line with direction of almost 0°). At HUA, ΔX versus ΔY points show no tilt on the regression line suggesting that the current system is almost towards east in this region as expected of Chapman [1] model. At ITI the ΔX versus ΔY points lie westward during the whole day suggesting a

northward tilt of the current vector in this region due to large declination.

In Fig. (2b) are shown the simultaneous developments of ΔX versus ΔZ at different stations HUA, ITI and TTB. At HUA the loop is symmetrical along the geographic meridian suggesting the location of HUA to be very close to the centre of EEJ current belt. The broadening of loop suggests that HUA was shifted slightly southward of EEJ centre in the forenoon and northward in the after noon hours. At ITI the direction of ΔX and ΔZ loop indicate that the station was throughout the day shifted slightly north of the EEJ centre. The ΔX versus ΔZ loop at TTB was very much tilted due to its proximity of the EEJ belt.

3.2. Daily Variation of ΔX , ΔY and ΔZ

In Fig. (3) are shown the daily variations of ΔX , ΔY and ΔZ for all the 37 days of observations respectively at the three stations (no data for 4 days at ITI). Corresponding variation of Dst index is also shown in the figure. Referring to Fig. (3c) the nighttime values of ΔX at any of the stations

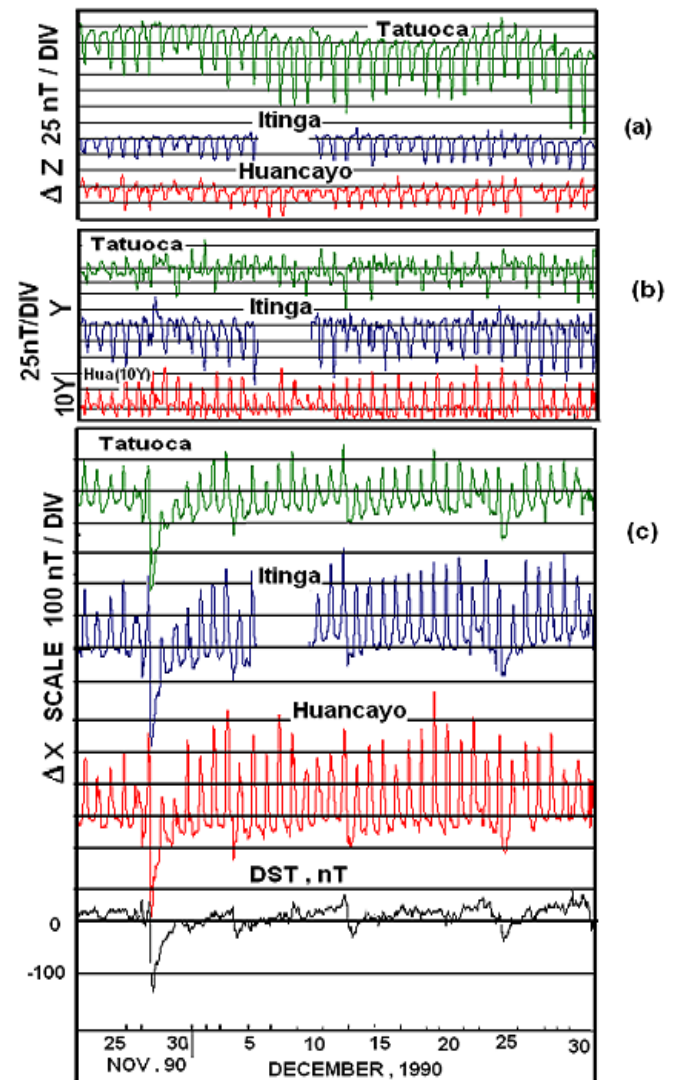


Fig. (3). Daily variations of (a) ΔZ , (b) ΔY and (c) ΔX for each day of the period 22 November-31 December 1990 at Tatuoca, Itinga and Huancayo. The variation of the Dst index is also shown.

show correspondence with the variation of Dst index. This is expected as there are no ionospheric currents during night time and the changes are primarily due to ring current. The signature of the magnetic storms can be seen in the figure especially for the one on 27 November 1990 is very clear. There are large day-to-day variations of the midday value of ΔX but definite associations between the different data sets are not very apparent.

Referring to Fig. (3b) showing the daily variations of ΔY at these stations, large positive deviation at HUA, large negative deviation at ITI and both positive and negative deviations at TTB are seen. No apparent effect of the magnetic storm is apparent at HUA but ΔY at ITI and TTB show positive deviations during the storm.

Referring to the daily variations of ΔZ in Fig. (3a) at these stations, comparatively large negative deviations of ΔZ are seen at TTB. No clear effect of Dst is seen.

3.3. Correlation Between Fluctuations (ΔX , ΔY) at the Three Stations

Next we checked correspondence between the midnight and midday values of X and Y at HUA and ITI. Deviations of X and Y were computed from overall mean value for midday and midnight hours at HUA and ITI. The mass plots of the deviations of X at HUA and ITI are shown in Fig. (4a). At midday, deviations in X at HUA are positively related to deviations in X at ITI with a correlation coefficient of 0.78. The regression line has a slope of 0.68. At midnight too, deviations in X at Huancayo are positively related to the deviations in X at ITI with correlation coefficient of 0.8 and the slope of regression line was 1.10. Thus the day-to-day changes of the ratio of X field at HUA to at ITI are 0.68 for the midday and 1.10 for midnight.

In Fig. (4b) are shown the mass plot of the deviations in Y field at HUA versus the same at ITI for the midnight and midday hours. The deviations in Y at HUA plotted are 10 times the actual values. During midnight deviations in Y at HUA are positively correlated with the deviations in Y at ITI with correlation coefficient of 0.72. During the daytime there is large scatter and no correspondences between the day-to-day changes of Y field at the two EEJ stations can be seen.

Correlations between the deviations of X and Y at HUA and ITI were computed for each of the hours. The daily variations of the correlation coefficient (blue lines) and of the slope of the regression line (pink line) are shown in Fig. (5a, b) respectively. The correlation values between HUA (X) versus ITI (X) range from about 0.94 to 0.54. The correlation coefficient is 0.90 or more from about 15 h to 02 h LT. It decreases to 0.54 around 0830 h LT and then rises slowly. The value of the slope is also lower between 0830 and 1430 h LT.

In Fig. (5b) are shown the daily variations of the correlation coefficient between HUA(Y) versus ITI (Y). It is to be noted that during midnight to morning (00-06 h LT) the values of the correlation coefficient as well as of the slope of the regression line are between 0.6 and 0.9. From 06 h LT the values of the correlation coefficient and the slope of the regression line decrease rapidly with values of 0.0 at 0930 h LT. The values increase then till 18 h LT to 0.50 for correlation coefficient and 0.36 for the slope of regression

line. The values show another sharp decrease to 0.0 at 2030 h LT and then increase to values of 0.50 at midnight. In general there is similarity in the daily variations of the correlation coefficients and slope of the regression lines between the deviations in X and Y at the two locations with minimum values around 0830-0930 h LT. Also the values are higher during night hours than day hours.

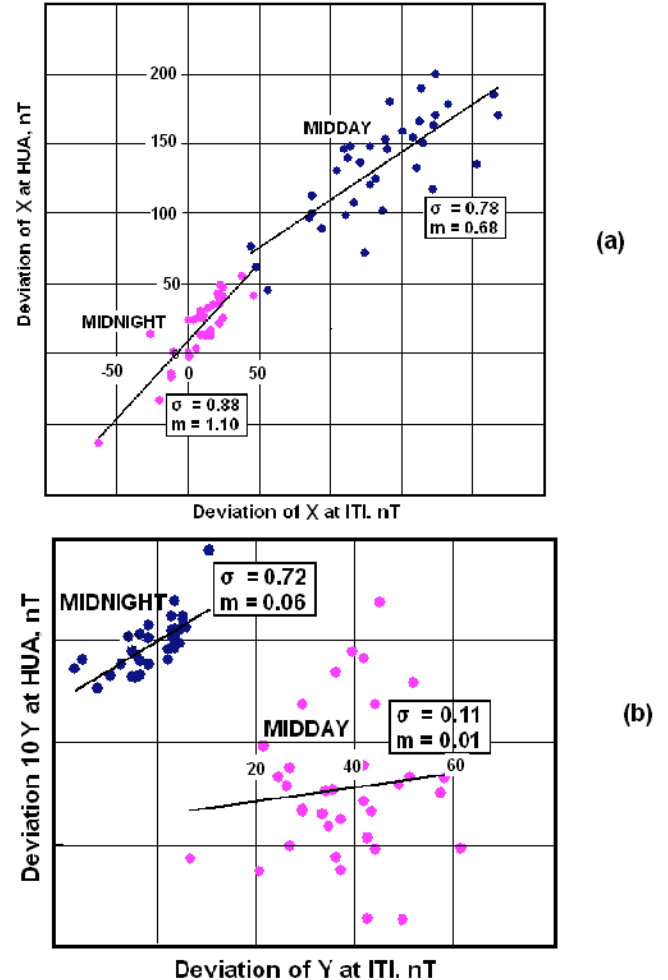


Fig. (4). (a) Scatter plots of the midday and midnight values of the deviations in X at HUA versus midday and midnight values of the deviations in X at ITI. The best-fit lines and the values of the standard deviations and slopes are also marked. (b). Scatter plots of the midday and midnight values of the deviations in Y (ten times) at HUA versus midday and midnight values of the deviations in Y at ITI. The best-fit lines and the values of the standard deviations and slopes are also marked.

3.4. Equatorial Electrojet and Disturbance Ring Current

A comparison of the plots of ΔX at HUA, ITI and TTB with the corresponding plots of Dst indicates a correlation between the equatorial ΔX and Dst as expected. James *et al.* [38] have shown significant correlation between day to day fluctuations of ΔH at the equatorial and mid latitude stations and corresponding Dst index for any time of the day.

In Fig. (6a) are shown the mass plots of X field at 00.0 hr LT (red dots) and 23.5 h LT (blue dots) versus Dst index for the three stations.

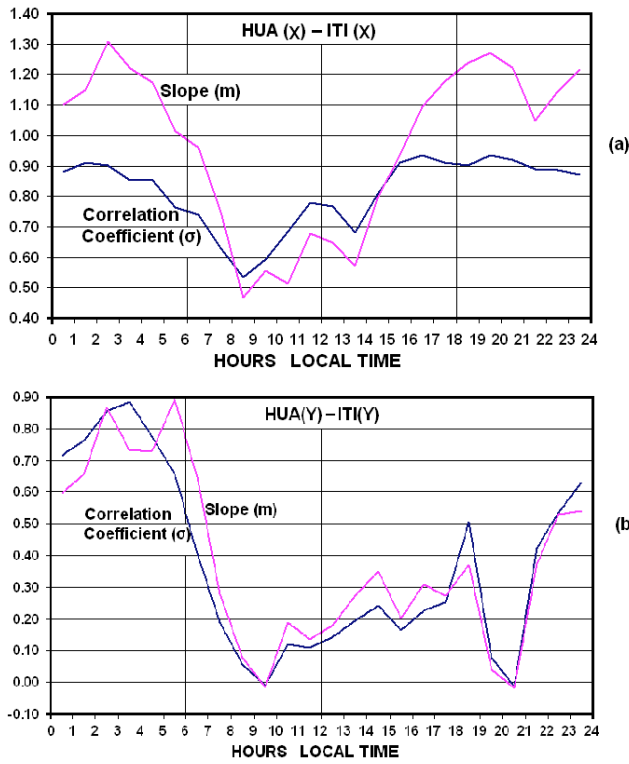


Fig. (5). The daily variations of the slopes of the best-fit line and the correlation coefficient between deviations in (a) X at HUA versus X at ITI and (b) Y at HUA versus Y at ITI.

There is a very high correlation between day-to-day fluctuations of night value of X field and Dst index with a correlation coefficient 0.9-1.0 and the slope of around 1.0 at any of these stations. Thus it can be concluded that the fluctuations in the night time values of X field at equatorial and low latitude stations are entirely due to the corresponding fluctuations of the ring current index. Uozumi *et al.* [18] from MAGDAS/CPMN data at Addis Ababa (AAB), Davao (DAV), Ancon (ANC) and Eusebio (EUS) reported that the mean night time (18-06 LT) value of the H showed variations like Dst.

Fig. (6b) shows similar mass plots of X at 11.5 hr LT (blue dots) and 12.5 hr LT (red dots) versus Dst index for the three stations. There are still positive correlations between midday X field and Dst index but the correlation coefficient values are low being 0.5 to 0.7.

The hourly mean Dst index were subtracted from the corresponding hourly mean values of X at the three stations to obtain a purer index of ionospheric contribution, X*,

$$X^* = X - \text{Dst} \tag{1}$$

The daily variations of X* at HUA, ITI and TTB during the period 23 November to 31 December 1990, are compared with corresponding Dst index in Fig. (7a). It is seen that the night time fluctuations in X (shown in Fig. 4c) are now gone, and the values of X* at any of the stations are fairly constant during the night time hours, in spite of the fluctuations of Dst during the same period.

In Fig. (7b, c) are shown the daily variations of the maximum (X*^{max}) and minimum (X*^{min}) values of X*, respectively during the entire period for the three stations. It

is very clear that the day-to-day variations of X* at any station are now very well correlated with each other. The fluctuations are again seen to be larger at Huancayo than at Itinga. The fluctuations of minimum value of X* at different stations are also seen to be well correlated.

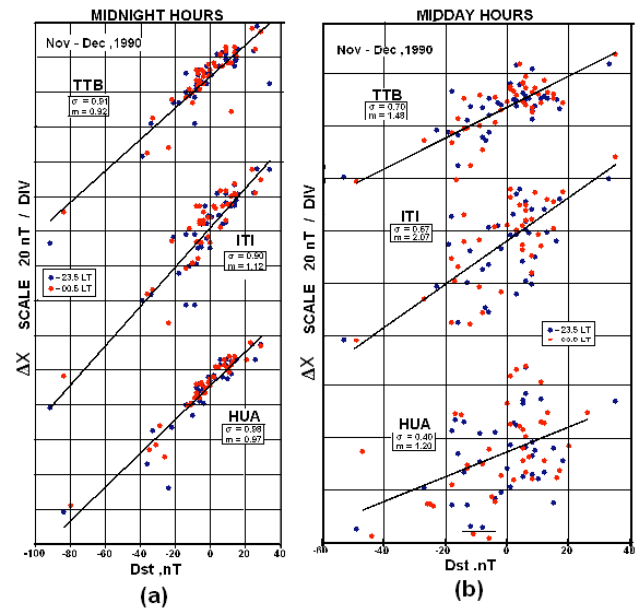


Fig. (6). Scatter plots of the deviations in X versus Dst index for TTB, ITI and HUA (a) at midnight hours and (b) at midday hours.

In Fig. (8) are shown the mass plots of hourly values of X* at HUA and ITI separately for the maximum daily values (X*^{max}) and for the minimum daily values of (X*^{min}) against the corresponding value at ITI. It is seen that X*^{max} at HUA are well correlated with X*^{max} at ITI with correlation coefficient of about 0.7. The mass plot of X*^{min} values at HUA and ITI almost overlap with good correlation of 0.80 between the two.

4. CONCLUSION

Characteristics of the equatorial electrojet at Huancayo (HUA) in western side of South America where the geomagnetic field is aligned almost along the geographic meridian and at Itinga (ITI) in eastern part of South America where the geomagnetic field is aligned about 19° west of the geographic meridian are studied. The mean intensity of the magnetic field in the two regions are almost of the same order. Comparisons are also made of the current at Itinga and at Tatuoca (TTB), a low latitude station in the same longitude sector.

The daily range of horizontal component of the geomagnetic field, H, is shown to be almost 16% higher at HUA compared to that at ITI. The daily variation of the eastward field, Y, showed a strong minimum of -40 nT around 13-14 hr LT at ITI whereas a positive peak of about 4 nT was observed at HUA around 11-12 hr LT. The vertical field, Z showed abnormally large negative values of -70 nT at TTB around 13 hr LT.

The day-to-day fluctuations of midday and midnight values of X field were positively correlated between HUA and ITI with correlation coefficient of 0.78 and 0.88

respectively. Values of Y field too were significantly positively correlated between HUA and ITI for midnight hours (0.72) while no correlation (0.11) was observed for the mid day hours.

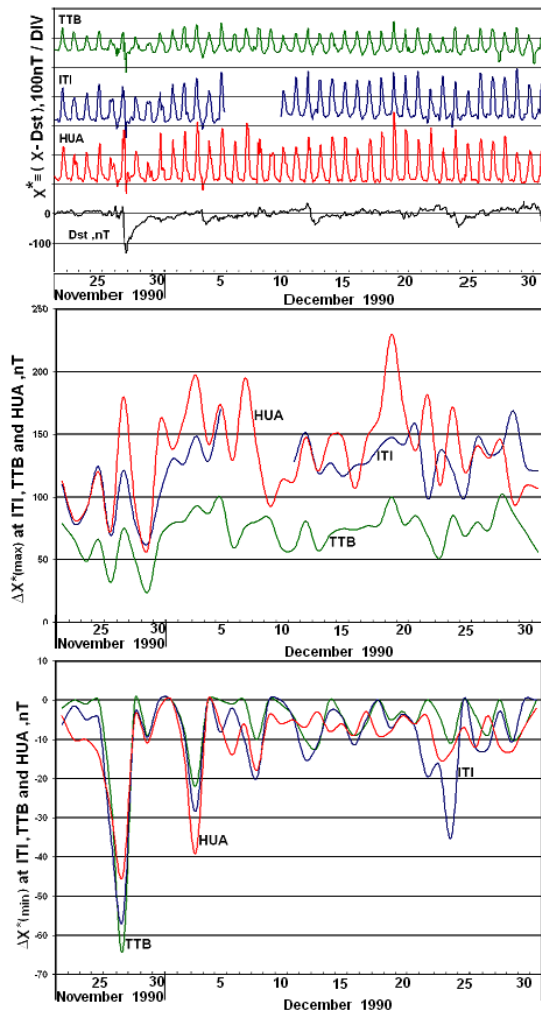


Fig. (7). (a) Daily variations of X^* , the deviations in X after subtracting the Dst values, for each of days during the period 22 November -31 December 1990 at TTB, ITI and HUA. The Dst values are also plotted. (b) Day-to-day variations of X^{*max} , the maximum value of the deviations in X after subtracting the Dst values, at HUA, ITI and TTB for the period 22 November-31 December 1990. (c) Day-to-day variations of X^{*min} , the minimum value of the deviations in X after subtracting the Dst values, at HUA, ITI and TTB for the period 22 November- 31 December 1990.

The midnight values of X field at HUA, ITI and TTB showed significant (0.90 or greater) correlation with Dst index. Correlation values of about 0.7 were observed between Dst and midday values of X at ITI and TTB and to a lesser degree (0.4) at HUA.

The ground geomagnetic field variations are assumed to be primarily due to the electric current in the ionosphere (E-region), which itself is the combined effects of the neutral wind and electrical conductivity (ionization density). Due to the interaction of the solar wind with magnetosphere, there is always dusk to dawn electric field imposed on the

ionosphere. In addition there are other currents namely magnetopause and magnetospheric tail currents affecting the ground magnetic measurements. Disturbance current which is always present even during the quiet days is additional source of magnetic field variation at ground. Thus there are complex sources of electric field imposed on the equatorial ionosphere besides the primary atmospheric dynamo electric field.

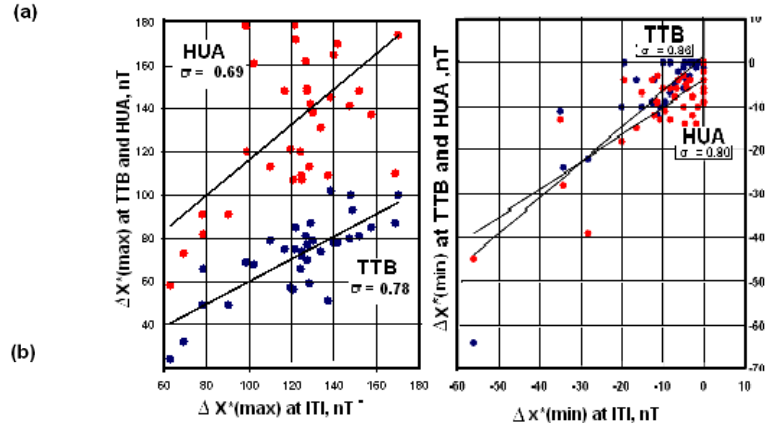


Fig. (8). Scatter plots of X^{*max} , the maximum deviations in X^* , at HUA and TTB plotted versus X^{*max} at ITI (left panel) and of X^{*min} , the minimum deviations in X^* , at HUA and TTB plotted versus X^{*min} at ITI (right panel).

CONFLICT OF INTEREST

The authors confirm that this article content has no conflict of interest.

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