Evaluation of Hydraulic Characteristics of the Normal Fault Based on Key Data

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Abstract: Based on the data of geological drill holes, roadway drifting, transient electromagnetic exploration, mine water inrushes, hydro-geochemistry and water levels on each side of the Haotang Normal Fault, this paper studies the water-bearing capacity and water conductivity of the fault in the Pingdingshan Tian’an Mining Company Limited No. 7 Coal Mine. The results indicate that the Haotang Fault is subject to compressive or shear-compressive stresses. The water-bearing capacity and water conductivity of the fault is poor. The water content of the Cambrian limestone in footwall is higher than that in hanging wall, and the mining threat of 15 coal seam in the south district (footwall) is greater than that in the north district (hanging wall).

Keywords: High-angle normal fault, Hydraulic characteristics, Mine water inrush, Water conductivity, Water content.

1. INTRODUCTION

Coal mine water inrush accidents are closely related to the complexity of geological conditions [1-3]. The depth of overburden and the stratigraphical relationships are usually varied and complex, particularly near the fault. This not only seriously affects the efficiency and safety of mining, but leads to significant changes in mine-water in flow. Therefore, a study of the spatial structure and hydraulic characteristics of the fault was conducted to identify changes in the flow of water into the mine and prevent water inrush.

The structure and hydraulic characteristics of the fault are predominantly studied using geological methods, hydrogeological methods, hydrogeochemical methods, geophysical prospecting, etc. However, because the spatial structure and hydraulic characteristics are affected by many factors, in practice only one or two methods are used [4-14].

In this study, using data of geological drill holes, roadway drifting, transient electromagnetic exploration (TEM), mine water inrush history, hydro-geochemistry, and the water levels on both sides of the Haotang Normal Fault [15-17], the water yield properties and water conductivity of the fault are analyzed, and their influences in the process of 15 coal seam mining are investigated.

2. LOCATION AND QUIFERS

The Pingdingshan Tian’an Mining Company Limited No.7 Coal Mine (No.7 Mine) is located at the north-west of the city of Pingdingshan, China, at 33°44’30"-33°45’25"N, 113°13’31.5"-113°16’56.5"E.

The mine is approximately 5.2 km long from east to west, and 1.6 km wide from north to south, covering an area of about 8.3 km². It is divided into north district and south district by the Haotang Fault (Fig. 1). The No. 7 Mine produces coal from the 15 coal seam at the base of the Carboniferous Pennsylvanian Formation (Fig. 2). The buried depth of 15 coal at south and north districts is 86-400 m and 400-523 m respectively.

As shown in Fig. (2), the immediate roof aquifer of the 15 coal is the L5 limestone of the Carboniferous Taiyuan Formation. The immediate floor aquifer is the L6+7 limestone aquifer overlying the Cambrian limestone aquifer (CL). The L5 limestone aquifer is thin (with the average thickness of 4.5 m), and is only weakly recharged with water, so its threat to the mining of 15 coal is relatively small.

The L6+7 limestone aquifer is thin (with the average thickness of 5 m), and the underlying aluminium mudstone has an average thickness of 3 m (with the minimum thickness 0.4 m). The CL aquifer is more than 300 m thick [18]. It is a confined aquifer and has an extremely uneven water yield. Because these are closely connected hydraulically and the water levels aquifer are generally identical in the L6+7 and the CL, both the L6+7 and CL pose severe threat due to the mining of 15 coal.

In the south district of the No. 7 Mine, the CL groundwater is recharged by atmospheric precipitation and river at the emergence zone of the CL (Fig. 1).

Since the recharged area is large and the flow distance is shorter in the CL groundwater, and the normal mine water
Fig. (1). Map of the No. 7 Mine showing the location of the Haotang Fault.

<table>
<thead>
<tr>
<th>System</th>
<th>Formation</th>
<th>Poles thickness</th>
<th>Lithologic column</th>
<th>Coal seam sign layer</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>Permain system</td>
<td></td>
<td></td>
<td>Purple mudstone</td>
</tr>
<tr>
<td></td>
<td>Shanxi Formation</td>
<td>12.90</td>
<td></td>
<td>K₁₈ Fragrant charcoal sandstone</td>
</tr>
<tr>
<td></td>
<td></td>
<td>5.00</td>
<td></td>
<td>K₁₇ Dazhan sandstone</td>
</tr>
<tr>
<td></td>
<td></td>
<td>4.85</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td></td>
<td>1.20</td>
<td></td>
<td>22 Coal</td>
</tr>
<tr>
<td></td>
<td></td>
<td>3.85</td>
<td></td>
<td>21 Coal</td>
</tr>
<tr>
<td></td>
<td></td>
<td>6.75</td>
<td></td>
<td>K₁₉ Mudstone</td>
</tr>
<tr>
<td></td>
<td>Carboniferous System</td>
<td>4.00</td>
<td></td>
<td>L₁ Limestone</td>
</tr>
<tr>
<td></td>
<td>Taiyuan Formation</td>
<td>4.50</td>
<td></td>
<td>L₅ Limestone</td>
</tr>
<tr>
<td></td>
<td></td>
<td>1.55</td>
<td></td>
<td>15 Coal</td>
</tr>
<tr>
<td></td>
<td>Base Formation</td>
<td>6.75</td>
<td></td>
<td>L₆⁺₇ Limestone</td>
</tr>
<tr>
<td></td>
<td>Guishan Formation</td>
<td>5.00</td>
<td></td>
<td>K₂₁ Purple mudstone</td>
</tr>
<tr>
<td></td>
<td></td>
<td>5.00</td>
<td></td>
<td>Cambrian limestone</td>
</tr>
</tbody>
</table>

Fig. (2). Columnar schematic diagram of the coal seam overlying the Cambrian limestone. Source: [15].
inflow is 2197 m³/h in the south district, which could be as high as 4750 m³/h. In the north district of the No. 7 Mine, the CL aquifer is far from the recharged area and cut by the Haotang fault, which made the normal mine water inflow dependent on the water yield properties and the hydraulic conductivity of the Haotang fault. The present study on the hydraulic characteristics of the Haotang Fault would be significant for establishing technical safety measures for mining the 1₃ coal.

3. FAULT DESCRIPTION AND WATER YIELD PROPERTIES

3.1. Fault Description

The Haotang Fault is concealed below the surface for about 200 m, along the east-west of No. 7 Mine (Fig. 1). It is a normal dip-slip fault striking at 120° and plunging at 30°. Its dip angle is 40-60°, with 10-35 m vertical displacement. The southern block of the fault has been uplifted relative to the northern block.

Fig. (3) shows three typical north-south cross-sections of the coal measures and aquifer, with the locations of the Haotang Fault and other minor faults. In section 1-1, the CL in the north and south districts are seen to be directly linked. Section 2-2 shows that the CL in the south district is directly connected to the 1₃ coal. Section 3-3 illustrated the connection between CL and the L₄₅ limestone in north district. The distinct strata connections at the different locations illustrate the complex influence of the fault on the various strata.

3.2. Fault Water Yield

3.2.1. Drilling and Roadway Engineering Details

Six drill holes (40-23, 39-20, 37-19, 35-29, 34-22 and 33-25) and nine roadway engineering sites in the Ji₁, Ji₂, and Ji₃ mining areas intersect the Haotang fault. The fracture plane become undulated gradually along the trend. Scars and steps obliquely distributed on hanging wall. Observation and analysis of the mechanical properties at these sites have shown that the fault is subjected to compressive or shear-compressive stresses.

Drilling and roadway engineering records show that the CL karst fissures are poorly developed in the hanging wall, the footwall, and the fractures in the Haotang Fault. No water was observed from any of the drill holes nor at any of the roadway drifting project sites intersecting the fault, illustrating the low water conductivity of the fault.

3.2.2. TEM Method

To survey the water yield properties of the fault, we carried out floor TEM on lines parallel to the fault along the western roadway. The No. 3 detecting line was 780 m long, beginning from the north of the Haotang Fault, passing through the fault to the southern side, then turning north westward and finally re-entering the fault zone (Fig. 4).

Fig. (5) (see the next page) is a cross-sectional diagram of the floor TEM results on detecting line No. 3. Clearly, the apparent resistivity was higher at 220-250 m, where the detecting line cut across the Haotang Fault and was highest at the finishing point (760-780 m) in the fault zone. Normally, higher apparent resistivity indicates poorly developed limestone karst fissures and associated low water yield. This is a further evidence that the water yield of the Haotang Fault is weak.

4. FAULT WATER CONDUCTIVITY

The water conductivity of the Haotang Fault was comprehensively and accurately determined from the water inrush when excavating the 1₃ coal, the chemical characteristics of the water in the CL, the water level distribution, and the TEM results.

4.1. Water Inrush Situation

The working face (22090, 22050, 22060, 22040) in the 1₃ coal was located near the Haotang Fault in the south district. The Gengæ haulage dip and track dip are located in the north district. The statistics of the frequency and volume of the mine water inrushes in 2002-2010 show that they occurred 53 times in the south district (34 times from the floor and 19 times from the roof; see Table 1). No inrushes occurred in the north district, where the total water inflow did not exceed 10 m³/h during roadway drive age, and the flow from the roof was only 1 m³/h. Clearly, floor water in the south district seriously threatens mining in the 1₃ coal, and the significant differences of floor water yield in north and south of the fault are obvious.

Fig. (3). Cross-sections through the Haotang fault. Source: [15].
Six underground water samples of the CL were collected from both sides of the Haotang Fault in August 2012 (Fig. 1). Sampling point 1 and 2 (No.1 point and No. 2 point) were in the north district, the others were in the south district. The proportions of K$^+$, Na$^+$, Ca$^{2+}$, Mg$^{2+}$, Cl$^-$, SO$_4^{2-}$, and HCO$_3^-$ ions, total dissolved solids and pH are listed in Table 2.

As shown in Table 2, north of the Haotang Fault, the proportions of ions were basically the same on sampling points 1 and 2, which indicates a single groundwater recharge path in the CL aquifer. South of the Haotang Fault, the proportions of ions were considerably different at different sampling points, which implies that the groundwater recharge path in the southern CL aquifer is complex and is strongly affected by various water sources.

The chemical constituents are totally different on both sides of the Haotang Fault, as seen in Table 2. Therefore, we inferred that the hydraulic connection between the north and south sides of the fault is particularly weak.

4.3. Water Level Distribution in the Aquifer

The water level in the CL was measured in September 2012. The observed values are shown in Fig. (6) as CL water table isolines on the southern side of the Haotang fault. It can

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**Table 1. Water inrush data in the south of the fault.**

<table>
<thead>
<tr>
<th>Type</th>
<th>Times</th>
<th>Quantity of Water Inrush</th>
<th>Remarks</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td></td>
<td>Max(m$^3$/h)</td>
<td>Average(m$^3$/h)</td>
</tr>
<tr>
<td>Floor</td>
<td>34</td>
<td>1350</td>
<td>78.6</td>
</tr>
<tr>
<td>Roof</td>
<td>19</td>
<td>450</td>
<td>23.6</td>
</tr>
</tbody>
</table>
be seen that the water table level is the lowest in the south of the Haotang fault, and increases with increasing distance southward. This indicates that, in the region of the No. 7 Mine, the aquifer is recharged by water entering at the southern outcrop of the CL.

On the northern side of the Haotang fault, the water levels are significantly different. The water table is the highest in the eastern area of the mine. This indicates that the water content of the CL aquifer is heterogeneous, and the hydraulic connection is weak.

The water levels are markedly different on either side of the Haotang fault, and the isogram is discontinuous, confirming the poor water conductivity of the fault.

### 4.4. TEM Method

Two floor TEM exploration lines perpendicular to the fault were established on the Geng2 track (Fig. 4). Line 1 (320 m long) was on the northern side of the Haotang fault, and line 2 (180 m long) was on the southern side of the fault.

#### 4.4.1. Line 1

Fig. (7) is the cross-sectional diagram of the vertical floor TEM on line 1. Apparent resistivity values less than 3 $\Omega \cdot m$ occurred in anomalous low-resistivity zones; the lowest apparent resistivity values were located at the Geng2 track dip roadway of 185-210 m and 260-285 m, and at depths of 55-105 m and 55-115 m. The rocks near the low resistivity zones are fractured karst or high water-content zones. None of the anomalies in the low-resistivity zone gave apparent resistivity readings less than 2 $\Omega \cdot m$, signifying the relatively small water-bearing capacity of the overall CL aquifer on the northern side of the Haotang fault.

#### 4.4.2. Line 2

At the time of the readings on line 2, interference was small and the background apparent resistivity was as high as 54 $\Omega \cdot m$. Values below 10 $\Omega \cdot m$ occurred only in relatively anomalous low-resistivity zones. Fig. (8) is a cross-sectional chart of the floor TEM results, showing the locations of the lower apparent resistivity values at the Geng2 track rise roadway of 18-40 m and 84-180 m, and at depths of 30-90 m and 24-90 m. Zones of apparent resistivity less than 2 $\Omega \cdot m$ in the south of the Haotang fault are corresponded to the presence of well-developed karst fissuring in the CL, from which we inferred that the anomalous low-resistivity zones indicated large volumes of water stored in the strata. Subsequently, inrushes of water from the floor did in fact occur when these zones were breached by excavation for roadways development, which confirmed the validity of our inference.

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**Table 2. CL water quality test results (mg/L).**

<table>
<thead>
<tr>
<th>District</th>
<th>Number</th>
<th>$K^+Na^+$</th>
<th>Ca$^{2+}$</th>
<th>Mg$^{2+}$</th>
<th>Cl$^-$</th>
<th>SO$_4^{2-}$</th>
<th>HCO$_3^-$</th>
<th>Total Dissolved Solids</th>
<th>pH</th>
</tr>
</thead>
<tbody>
<tr>
<td>North</td>
<td>1</td>
<td>37.7</td>
<td>85.0</td>
<td>39.4</td>
<td>36.5</td>
<td>89.8</td>
<td>379.5</td>
<td>478.2</td>
<td>7.4</td>
</tr>
<tr>
<td></td>
<td>2</td>
<td>36.3</td>
<td>77.6</td>
<td>37.9</td>
<td>36.5</td>
<td>65.8</td>
<td>376.5</td>
<td>442.4</td>
<td>7.3</td>
</tr>
<tr>
<td></td>
<td>Average</td>
<td>37.0</td>
<td>81.3</td>
<td>38.7</td>
<td>36.5</td>
<td>77.8</td>
<td>378.0</td>
<td>460.3</td>
<td>7.4</td>
</tr>
<tr>
<td>South</td>
<td>3</td>
<td>20.9</td>
<td>100.0</td>
<td>9.1</td>
<td>17.4</td>
<td>120.1</td>
<td>223.1</td>
<td>379.0</td>
<td>7.7</td>
</tr>
<tr>
<td></td>
<td>4</td>
<td>58.4</td>
<td>72.5</td>
<td>42.4</td>
<td>50.0</td>
<td>101.8</td>
<td>373.4</td>
<td>511.9</td>
<td>7.3</td>
</tr>
<tr>
<td></td>
<td>5</td>
<td>32.2</td>
<td>122.4</td>
<td>25.8</td>
<td>30.1</td>
<td>167.6</td>
<td>311.8</td>
<td>534.4</td>
<td>7.3</td>
</tr>
<tr>
<td></td>
<td>6</td>
<td>20.7</td>
<td>85.0</td>
<td>31.8</td>
<td>29.1</td>
<td>41.8</td>
<td>370.4</td>
<td>393.6</td>
<td>7.3</td>
</tr>
<tr>
<td></td>
<td>Average</td>
<td>33.1</td>
<td>95.0</td>
<td>27.3</td>
<td>31.7</td>
<td>107.8</td>
<td>320.0</td>
<td>454.7</td>
<td>7.4</td>
</tr>
</tbody>
</table>

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**Fig. (6).** Isoline map of water table levels in the CL aquifer. **Source:** [15].
No apparent resistivity readings less than 2 Ω·m were observed on line 1, and the water-bearing anomalies covered a smaller area than those on line 2. These facts illustrate that the aquifer contains more water in the south of the fault than in the north of the fault. They also indirectly confirm that the CL aquifer groundwater contact isn’t close to both sides of the fault.

5. DISCUSSION

The above study indicates that the water-bearing capacity and water conductivity of the Haotang fault is poor. Since the CL water with great water-bearing capacity and water conductivity in south district couldn’t cross the fault to north district. These make the weaker water-bearing capacity and smaller mine water inflow in the north compared with the south. It is because the distance from 1_5 coal mine floor to the CL is only 16.75 m and the water pressure of 1_5 coal mine is about 0.8 MPa, the coal mining may easily lead to the water inrush. Therefore, in the process of 1_5 coal mining, we use geophysical exploration to confirm the CL water anomaly area, build the drill holes to get the water pressure, thickness of aquiclude and date of water-bearing capacity. These will help formulate the measure of water drainage hole, floor grouting, and water transport system to make sure the safe mining of 1_5 coal seam.

CONCLUSION

(1) Geological drill holes, roadway engineering and floor TEM data indicate that the Haotang fault is subject to compressive or shear-compressive stresses. No karst fissures in the CL are developed in the fault zone in either the hanging wall or the footwall, as a result of which it has a relatively low water content.

(2) In zones remote from the Haotang fault, the frequency of water inrushes, the hydrogeochemistry, the water level distribution, and the water content are all markedly different on either side of the fault, indicating that the CL hydraulic connection is not close and that water conductivity of the fault is particularly weak.

(3) Because the water-bearing capacity and water conductivity of Haotang fault is poor, the CL water with great water-bearing capacity and water conductivity in south district couldn’t cross the Haotang fault to north district. These make the weaker water-bearing capacity and smaller mine water inflow in the north compared with the south. While, in the north district of the fault, it presents a smaller threat to mining in the 1_5 coal seam.

CONFLICT OF INTEREST

The authors confirm that this article content has no conflict of interest.

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